#### Elements for the production of Objective Seasonal Forecasts: MedCOF sub-region

### ENSO as a source of predictability for the North Atlantic-Mediterranean region

**Bianca Mezzina** Université Catholique de Louvain, Earth and Life Institute bianca.mezzina@uclouvain.be

#### **ENSO BASICS: WHAT IS ENSO?**

#### ENSO = El Niño-Southern Oscillation



Recurring natural phenomenon in the tropical Pacific region



Coupled oscillation: variations in ocean (El Niño) and atmosphere (Southern Oscillation)



Fluctuation between 3 phases: Neutral, El Niño (warm) and La Niña (cold)



#### **ENSO BASICS: SEA SURFACE TEMPERATURE (SST)**

(a) Neutral



(a) DJF observed SST average during neutral years (b,c) Composite of DJF observed SST anomalies during El Niño and La Niña years

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#### **ENSO BASICS: WHEN AND WHY**



Standard deviation of observed monthly Niño3.4 index

#### **ENSO BASICS: LOCAL ATMOSPHERIC CIRCULATION**



#### **ENSO BASICS: LOCAL ATMOSPHERIC CIRCULATION**



#### **ENSO BASICS: LOCAL ATMOSPHERIC CIRCULATION**



#### **ENSO-NAE TELECONNECTION**



#### ENSO is a potential **source of seasonal predictability** for mid-latitude regions

### **ENSO-NAE TELECONNECTION**



ENSO is a potential **source of seasonal predictability** for mid-latitude regions

...And a very powerful one!

### **ENSO PREDICTABILITY**

ENSO can be predicted several **months in advance** 



**Figure 10.3** The forecast lead-time (*y*-axis) when Niño-3.4 index skill exceeds certain correlation thresholds. On the *x*-axis, the target month of the forecast is presented. Model data is based on the ensemble average of ~100 members from the North American Multi-Model Ensemble from 1982 to 2018. Departures in the Niño-3.4 index are based on 1982–2010 monthly averages.

### **ENSO-NAE TELECONNECTION**



El Niño and rainfall





No impact?

### **ENSO-NAE TELECONNECTION**



**El Niño and rainfall** 



Some impact, but hard to detect...



Brönnimann (2007)



Brönnimann (2007)

### ENSO-NAE TELECONNECTION: "CANONICAL" DIPOLE

Late-winter sea-level pressure (SLP) dipole



*Dipole* between **high latitudes** and **mid-latitude North Atlantic** 



*Robust* and *significant* in **late winter** (January-March; JFM)



Observed SLP anomalies associated with El Niño

#### **ENSO-NAE TELECONNECTION: "CANONICAL" DIPOLE**

Late-winter **sea-level pressure** (SLP) dipole



What about **Europe** and the **Mediterranean**?



Observed SLP anomalies associated with El Niño

#### ENSO-NAE TELECONNECTION: "CANONICAL" SIGNAL





Observed SLP anomalies associated with El Niño

Brönnimann (2007)

#### LATE-WINTER SIGNAL OVER EUROPE



An example from CMIP5 historical runs...



Regression of Nino3.4 index onto T850 and precipitation for two versions of the MPI-ESM model (historical runs, 1 member)

**Challenges** in detecting an ENSO signal in Europe:

• Large internal variability (low signal-to-noise ratio)

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- Large internal variability (low signal-to-noise ratio)
- Sampling issues:
  - Limited observational record (100 years)  $\rightarrow$  "few" ENSO events
  - In models: how many years/members do we need?

**Challenges** in detecting an ENSO signal in Europe:

- Large internal variability (low signal-to-noise ratio)
- Sampling issues:
  - Limited observational record (100 years)  $\rightarrow$  "few" ENSO events
  - In models: how many years/members do we need?
- ENSO diversity
- Non-stationarity? (more on this later!)











15°

15°N

15°W

0°

15°E

30°E

45°E

15°N •

15°W

0°

15°E

30°E

45°E

**EN** (a) EN SST 180 120°W

Benassi et al. (2021)

25

#### **1.** Via the troposphere

- 2. Via the stratosphere
- 3. Others

Quasi-stationary tropospheric Rossby wave train



Observed 200-hPa geopotential height anomalies associated with ENSO

Quasi-stationary tropospheric Rossby wave train

#### For a **free** wave (horizontal motion):



#### Quasi-stationary tropospheric Rossby wave train

#### For a **forced** wave (horizontal motion):





Observed 200-hPa geopotential height anomalies associated with ENSO













![](_page_32_Figure_1.jpeg)

Convection ↔ low-level moisture convergence ↔ SST gradient

![](_page_32_Figure_3.jpeg)

![](_page_32_Figure_4.jpeg)

![](_page_32_Figure_5.jpeg)

![](_page_32_Figure_6.jpeg)

Zonal gradient of total SST

**EN** 

EC-EARTH

**CNRM** 

CMCC

![](_page_33_Figure_2.jpeg)

![](_page_34_Figure_1.jpeg)

![](_page_35_Figure_1.jpeg)

#### Quasi-stationary tropospheric Rossby wave train

#### For a **forced** wave (horizontal motion):

![](_page_36_Figure_3.jpeg)

![](_page_36_Figure_4.jpeg)

Observed 200-hPa geopotential height anomalies associated with ENSO

#### Quasi-stationary tropospheric Rossby wave train

#### For **ENSO**:

![](_page_37_Figure_3.jpeg)

![](_page_37_Figure_4.jpeg)

Observed 200-hPa geopotential height anomalies associated with ENSO

Sardeshmukh and Hoskins (1988) Qin and Robinson (1993)

**EN** 

120°W

60°W

120°E

180°

30°5

120°E

![](_page_38_Figure_2.jpeg)

120°W

60°W

IN

EN

![](_page_39_Figure_2.jpeg)

![](_page_39_Figure_3.jpeg)

LN

![](_page_39_Figure_4.jpeg)

![](_page_39_Figure_5.jpeg)

![](_page_39_Figure_6.jpeg)

![](_page_39_Figure_7.jpeg)

![](_page_39_Figure_8.jpeg)

![](_page_39_Figure_9.jpeg)

EN LN EN (EC-EARTH) LN (EC-EARTH) TRWS200, Z200, v<sub>div</sub>200 TRWS200, Z200, v<sub>div</sub>200 Amplitude EN > LN **EC-EARTH** 60°I ------18 30°S Westward **shift** LN wrt to EN EN (CNRM) LN (CNRM)  $\mathsf{TRWS}200,\mathsf{Z}200,\vec{\mathsf{v}}_{\mathsf{div}}200$ TRWS200, Z200, v<sub>div</sub>200 **CNRM** 60°N 0' -18 30°S EN (CMCC) LN (CMCC)  $\mathsf{TRWS}200,\mathsf{Z}200,\vec{\mathsf{v}}_{\mathsf{div}}200$  $\mathsf{TRWS}200,\mathsf{Z}200,\vec{\mathsf{v}}_{\mathsf{div}}200$ 60°I CMCC 0' -18 30°S 120°E 180° 120°W 60°W 120°E 180° 120°W 60°W

#### **CHALLENGES**

**Challenges** in representing the ENSO-NAE teleconnection:

- 0. Getting "ENSO" right (SST pattern, timing, amplitude, ...)
- 1. Getting the right tropical response
- 2. Importance of the **mean flow**

#### **CHALLENGES**

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![](_page_42_Figure_5.jpeg)

- 1. Via the troposphere
- 2. Via the stratosphere
- 3. Others

# **ENSO impacts** on the winter **mean state** of the *polar stratosphere* have been shown

![](_page_44_Picture_2.jpeg)

# **ENSO impacts** on the winter **mean state** of the *polar stratosphere* have been shown

This stratospheric ENSO signal could *propagate downwards* and affect the surface (Stratospheric Pathway hypothesis)

![](_page_45_Picture_3.jpeg)

#### **STRATOSPHERIC PATHWAY**

![](_page_46_Figure_1.jpeg)

#### **STRATOSPHERIC PATHWAY**

**Downward** propagation of stratospheric signal

![](_page_47_Figure_2.jpeg)

Ineson and Scaife (2009)

![](_page_48_Figure_1.jpeg)

![](_page_49_Picture_1.jpeg)

![](_page_49_Picture_2.jpeg)

![](_page_49_Figure_3.jpeg)

![](_page_50_Figure_1.jpeg)

sure (hPa)

Pressure (hPa)

10

20

30

50

20° N

![](_page_51_Picture_1.jpeg)

What we expect

![](_page_51_Picture_2.jpeg)

°C

3.5 3.0 2.5 2.0 1.5

1.0

-0.5

-1.0

-1.5

-2.0

2.5

-3.0 -3.5 -4.0

2.5

-0.5

-1.0

-1.5

-2.0

2.5

-3.0

52

![](_page_51_Figure_3.jpeg)

**EC-EARTH** 

**CNRM** 

#### CHALLENGES

**Challenges** in representing the ENSO-NAE teleconnection:

- 0. Getting "ENSO" right (SST pattern, timing, amplitude, ... )
- 1. Getting the right tropical response
- 2. Importance of the **mean flow**
- 3. Having the expected **stratospheric signal** ( $\neq$  high resolution in the stratosphere) \*

#### **EXTRA-TROPICAL RESPONSE: SLP**

**EN** 

![](_page_53_Figure_2.jpeg)

EC-EARTH

**CNRM** 

CMCC

![](_page_53_Figure_3.jpeg)

30°E

45°E

\* Even though the contribution of the stratosphere to the surface JFM signal is not clear

![](_page_54_Figure_1.jpeg)

2. Via the stratosphere

![](_page_54_Figure_3.jpeg)

#### 3. Others

e.g.

- Delayed impacts through other basins (North Atlantic?)
- Secondary wave sources
- Splitting of main wave train

- ...

60

#### CHALLENGES

**Challenges** in representing the ENSO-NAE teleconnection:

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- 1. Getting the right tropical response
- 2. Importance of the **mean flow**
- 3. Having the expected **stratospheric signal** (≠ high resolution in the stratosphere)
- 4. Timing is everything

#### **INTRA-SEASONAL EFFECTS**

![](_page_56_Figure_1.jpeg)

### **INTRA-SEASONAL EFFECTS**

![](_page_57_Figure_1.jpeg)

### SEASONAL MODULATION

Precipitation

![](_page_58_Figure_2.jpeg)

![](_page_58_Figure_3.jpeg)

Shaman et al. (2006)

#### CHALLENGES

**Challenges** in representing the ENSO-NAE teleconnection:

- 0. Getting "ENSO" right (SST pattern, seasonal cycle, amplitude, ...)
- 1. Getting the right tropical response
- 2. Importance of the **mean flow**
- 3. Having the expected **stratospheric signal** (≠ high resolution in the stratosphere)
- 4. Timing is everything
- 5. Modulation by **other factors** (e.g. PDO)

### A QUICK LOOK AT THE PDO

![](_page_60_Figure_1.jpeg)

#### **PDO = P**acific **D**ecadal **O**scillation

![](_page_60_Picture_3.jpeg)

Leading mode of variability in the North Pacific

![](_page_60_Picture_5.jpeg)

**Low-frequency** variability (10-30 years)

![](_page_60_Picture_7.jpeg)

**Positive** and **negative** phases

![](_page_60_Picture_9.jpeg)

Combination of different processes

## EL NIÑO AND THE PDO

# Thank you MEDSCOPE

![](_page_61_Figure_2.jpeg)

## EL NIÑO AND THE PDO

![](_page_62_Figure_1.jpeg)

The PDO can modulate the **amplitude** of the ENSO-NAE teleconnection.

**PDO-** : amplification of canonical EN pattern

By **altering the mean flow** and favoring a more poleward propagation of planetary waves across the North Pacific-American sector

#### CHALLENGES

**Challenges** in representing the ENSO-NAE teleconnection:

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